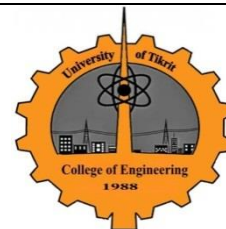


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Rainfall Data Analysis and Study of Meteorological Draught in Iraq for the Period 1970-2010

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Abstract

This work aimed to analyze and evaluate the meteorological draught in Iraq. Monthly rainfall data were collected from 22 meteorological stations scattered all over Iraq for the period 1970-2010. Various statistical tests have been performed to examine the data accuracy, e.g. Consistency test, Trend analysis, as well as homogeneity test. The Standard Precipitation Index (SPI) and Theory of Runs are used to analyze the meteorological draught and evaluating the draught characteristics. SPI values have been estimated for 12-month time scale, from January to December. To analyze the drought characteristics, the accumulative deficit, drought duration and intensity for each station have been estimated. The results showed that the rainfall data series considered in the study are all consistent with negative trends except Hilla station, random except Haditha and Samawah stations and homogeneous. It was shown that the study area has been suffered from sequent drought events through nearly half of the years considered in the study, and the worst were in 1997-2001 and 2007-2010 in which extreme droughts were dominated several parts of study area.

Keywords: Meteorological draught, Iraq, theory of runs, standard precipitation index.

تحليل بيانات الامطار ودراسة الجفاف المناخي في العراق للفترة 1970-2010

الخلاصة

تهدف هذه الدراسة الى تحليل وتقييم الجفاف المناخي في العراق. تم جمع بيانات الامطار الشهرية من 22 محطة انواء جوية منتشرة في كل انحاء العراق للفترة 1970-2010. تم اجراء العديد من الفحوصات الإحصائية لاختبار مدى دقة البيانات وموثوقيتها، منها فحص الاتساق وفحص الاتجاه العام فضلا عن فحص التجانس. استخدم في هذا البحث دليل المطر القياسي SPI ونظرية الحدث المستمر Theory of runs لتحليل الجفاف المناخي وتقييم خصائص الجفاف. تم حساب قيم SPI بمقياس زمني مقداره 12 شهر تبدأ من شهر كانون الثاني وتنتهي بشهر كانون الاول. لتحليل خصائص الجفاف تم تخمين قيم النقص المتراكم في قيم الامطار وطول فترة الجفاف وكثافة الجفاف ولكل محطة من المحطات المشمولة بالدراسة. أظهرت النتائج ان بيانات الامطار المعتمدة في البحث جميعها متجانسة ومتسقة مع وجود اتجاه عام سالب لبيانات جميع المحطات عدا محطة الحلة. كما كانت البيانات عشوائية عدا محطات حديثة والسماوة. بينت الدراسة ان منطقة الدراسة قد عانت من فترات جفاف متعاقبة حيث وصل عدد السنوات التي عانت من الجفاف الى نصف سنوات فترة الدراسة وكانت اسوأ فترات الجفاف هي 1997-2001 و 2007-2010 حيث ساد فيهما الجفاف المتطرف وفي كل أجزاء منطقة الدراسة.

الكلمات الدالة: الجفاف المناخي، العراق، نظرية الحدث المستمر، دليل المطر القياسي.

Introduction

Drought is a temporary feature resulting from prolonged absence, or deficiency or poor distribution, of precipitation (Ogallo, 1994)[1]. It is a natural recurrent phenomenon, which occurs in different temporal and spatial scales. Drought is difficult to define due to its strong dependency on time and space, and its variety of effects.

A drought is an indistinct event of water deficiency that results from the combination of many integrated factors, and neither the start nor the termination can be exactly defined (Kossida et al. 2009)[2]. Although numerous interpretations of drought have been offered, the most significant determinant of drought is the amount of precipitation an area gets compared to normal (Edwards and Mckee, 1997)[3]. Types of drought can be defined by the effect on specific scopes and are classified into meteorological, agricultural, hydrological and socio-economic droughts. Meteorological drought is associated with a precipitation deficit and it depends upon its duration, which can cause agricultural (related to soil moisture) or hydrological drought (related to e.g. stream flow, groundwater level, or reservoir storage). Socio-economic drought addresses the monetary effects of drought (Norouzi et al., 2012)[4].

Drought planners usually rely on some mathematical indices to decide when to carry out water conservation or mitigation measures against drought (Loucks and Bee 2005)[5]. Drought indices are tools for measuring and analysing drought severity. These indices reflect the impact of drought on different types of water sources. Such indices differ by drought type and variables included in these indices (Wilhite and Glantz 1985, Svoboda and Coauthors 2002)[6,7]. Widely used indices in drought studies include the Palmer Drought Severity Index, PDSI (Palmer, 1965)[8], the Standardized Precipitation Index, SPI (McKee et al. 1993)[9], Effective Drought Index, EDI (Byun and Wilhite, 1996)[10], Reconnaissance Drought Index, RDI (Tsakiris et al. 2007)[11], Standardized Precipitation Evapotranspiration Index, SPEI (Vicente-Serrano et al., 2010)[12], Soil Moisture Index, SMI (Nam

et. al. 2012) [13], Integrated Surface Drought Index, ISDI (Wu et. al. 2013) [14], Multivariate Standardized Drought Index, MSDI (Hao and AghaKouchak, 2013) [15]. Although that the literature showed verity of drought indices, but the Standard precipitation Index SPI is the most used index (Edossa et al. 2010)[16].

Iraq is one of the countries in the middle-east region, which suffered from frequent drought events in the last decades. This had serious influences on the water resources, irrigated and rainfed agriculture, areas of cultivated lands, desertification and demographic distribution, which consequently caused serious economic problems for the country. The research works regarding droughts in the country are rare, therefore, the current work aims to study and analyse the meteorological drought in Iraq using the SPI. The study included analysing the rainfall data with different methods to ensure its homogeneity, consistency and randomness.

Study Area

The study considered the entire area of Iraq within its borders. Iraq lies in a semi-arid region between longitudes 38° 45' and 48° 45' and latitudes 29° 5' and 37° 22' with an area of 437,049 km². It has different terrain types including mountainous territory in the north and north-east, desert territories in the west and south-west, and marshlands in the south, resulting in different climate characteristics from a region to another. Monthly rainfall data have been collected from 22 meteorological stations (i.e. Mosul, Duhok, Erbil, Sulaymaniyah, Kirkuk, Khanqin, Tel-Afer, Dukan, Darbandikhan, Baiji, Baghdad, Qaim, Rutba, Hai, Nasiriya, Hilla, Basra, Nukheb, Ramadi, Samawa, Hadithah and Sinjar) for the period 1970-2010. Figure (1) demonstrates the study area and locations of meteorological stations considered in the study, and Figure (2) shows the spatial distribution of the average annual rainfall for the period 1970-2010.

Figure (2) proves that the annual rates of rainfall varies with the terrain. Highest rates of annual rainfall are noticed in the north-eastern part of Iraq, including Dukan, Sulymania, and Darbandikhan, which range from 659 to 736 mm. Rates of annual rainfall decline on southwards and westwards

direction. Lowest average annual rainfall rates have been recorded in Nukheh station in the south-western part of Iraq with only 74 mm.

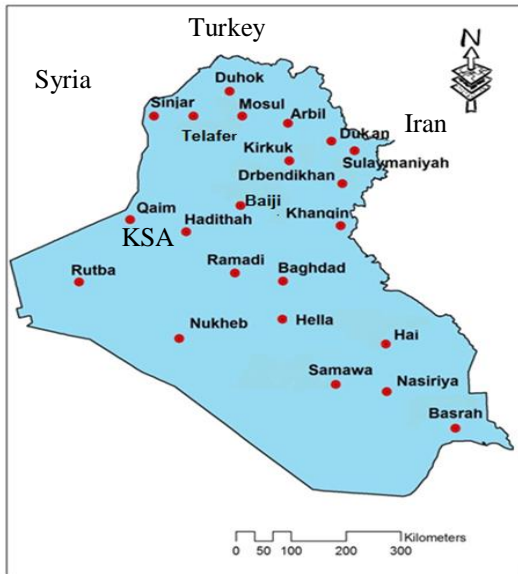


Fig. 1. The study area and locations of the selected meteorological stations

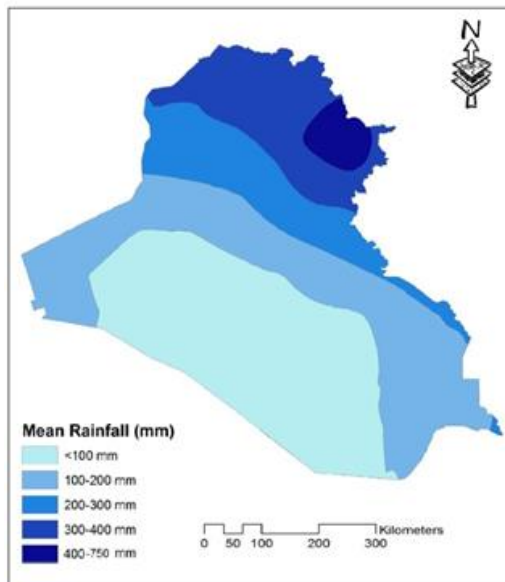


Fig. 2. The spatial distribution of annual rainfall for the period 1970-2010

Data Preparation

Precipitation is considered as the main natural source of fresh water resources,

especially for arid and semi-arid environments. However, the rainfall at these environments characterized with low scale depth, non-uniform distribution with inequality quantities from year to year. Statistical tests, annual precipitation accumulates and their spatial and temporal distribution estimation are essential procedures for the data collected from meteorological stations prior to any draught study. Table (1) shows the salient statistical features of rainfall data of the selected stations.

Estimation of Missing Rainfall Data

Rainfall records in all meteorological stations covered by the study experience different periods of missing data which may reach in some stations as long as twelve years, continuous or intermittent, and differ from one station to another. Thus, missing data values have been estimated by adopting two methods, i.e., Arithmetic Average Method and Normal Ratio Method. These two methods rely on the estimate of specific station precipitation values by considering the precipitation data of neighbouring stations (Varshney et al., 1979)[17].

The records of two stations, i.e., Ramadi and Qaim stations, have a lack of 12 years. It is difficult to estimate these data by the above-mentioned methods, thus, the linear regression method was used to link the data of the aforesaid two stations and the neighbouring stations. The following equation has been suggested to estimate the missing data for Ramadi station using the data of Ramadi, Rutba, Baiji, Haditha, and Baghdad stations with regression coefficient (R^2) of 0.709.

$$R_{Ramadi} = 0.444 + 0.0823 R_{Haditha} + 0.144 R_{Baghdad} + 0.259 R_{Baiji} + 0.29 R_{Rutba} \dots \dots \dots (1)$$

As for Qaim station, the following equation has been suggested with a regression coefficient (R^2) of 0.588.

$$R_{Qaim} = 3.667 + 0.725 * R_{Rutba} \dots \dots \dots (2)$$

Where R_x is the monthly rainfall data (mm) at station x

Table 1. Preliminary statistics of rainfall data for the selected meteorological stations

Station	Elevation (m)	Longitude	Latitude	Annual rainfall mean (mm)	Standard deviation (mm)	Minimum annual rainfall (mm)	Maximum annual rainfall (mm)
Nukheb	305	42° 15'	32° 02'	74	38	24	200
Samawa	6	45° 16'	31° 18'	92	50	15	228
Hilla	25	44° 26'	32° 11'	94	36	41	192
Rutba	615.5	40° 17'	33° 02'	114	59	23	264
Ramadi	48	43° 19'	33° 27'	114	46	56	241
Baghdad	32	44° 14'	33° 14'	117	49	50	284
Nasiriya	3	46° 14'	31° 05'	121	53	46	246
Qaim	178	41° 10'	34° 23'	121	58	23	284
Hadithah	140	42° 22'	34° 04'	122	58	39	274
Hai	14.9	46° 03'	32° 10'	134	57	55	290
Basra	2.4	47° 43'	30° 34'	136	57	32	297
Baiji	115	43° 29'	34° 55'	198	71	85	377
Khanqin	202.2	45° 26'	34° 18'	289	88	72	457
Tel-Afer	200	42° 29'	36° 22'	314	112	134	614
Kirkuk	330.8	44° 24'	35° 28'	355	132	135	696
Sinjar	538	41° 50'	36° 19'	355	134	164	663
Mosul	223	43° 09'	36° 19'	358	111	194	633
Erbil	426	44° 20'	36° 11'	420	128	157	886
Duhok	536	43° 00'	36° 50'	553	173	277	911
Darbandikhan	400	45° 45'	35° 08'	659	186	296	1041
Sulaymaniyah	853	45° 27'	35° 33'	706	178	339	1085
Dukan	490	44° 57'	35° 57'	735	221	366	1468

Theoretical Background

Statistical Tests of Rainfall Data

The statistical analyses of hydrological time series data considered in water resources planning studies are based on a set of major assumptions, e.g. the series is consistent, is trend-free and makes up a stochastic process whose random part follows the appropriate probability distribution function. Consistency indicates that all the collected data belong to the same statistical population. Trend exists in a data set if there is a significant positive or negative correlation between time and the observations. Trend is normally presented through human activities. Randomness in a hydrological time series means that the data arise from natural causes. If there is no randomness, then the series is persistent; this persistence is normally quantified in terms of the serial correlation coefficient (Adeloye and Montaseri 2002)[18]. In this paper, the rainfall data series of the selected stations have been encountered several statistical analysis to prove their consistency, randomness and homogeneity.

Test for Consistency of record

Double mass curve is used to check the consistency of many kinds of hydrologic data by comparing data for a single station with that of a pattern composed of the data from several other stations in the area. The double-mass curve can be used to adjust inconsistent precipitation data. The graph of the cumulative data of one variable versus the cumulative data of a related variable is a straight line so long as the relation between the variables is a fixed ratio. Breaks in the double-mass curve of such variables are caused by changes in the relation between the variables. These changes may be due to changes in the method of data collection or to physical changes that affect the relation (Adeloye and Montaseri, 2002)[18]. Accumulated rainfall for the specific station as well as the accumulated values for adjacent main stations shall be calculated. Then, the relation is drawn between them as shown in Figure (3).

$$P_{cx} = P_x * \frac{M_c}{M_a} \dots\dots\dots (3)$$

Where P_{cx} is the corrected precipitation at time period t_1 at station x , P_x is original recorded precipitation at time period t_1 at station x , M_c refers to the corrected slope of the double mass curve and M_a is the original slope of the mass curve.

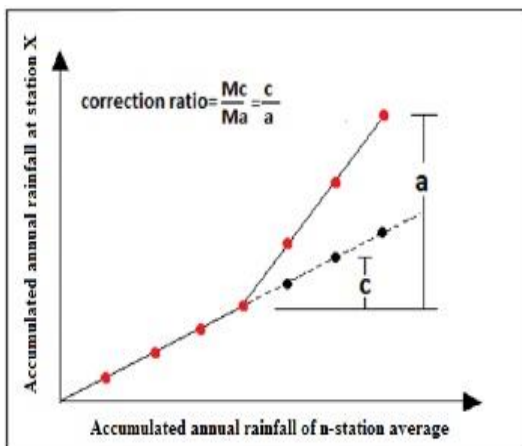


Fig. 3. Annual rainfall data consistency test

Precipitation Trend Analysis

Trend is defined as the slight change of a phenomenon from time to another with no impact except after a long time. When numbers indicate to increase, the direction is positive, while the opposite is negative. Many statistical techniques (parametric or non-parametric) have been developed to detect trends within time series such as linear regression method, Spearman's Rho test, Mann-Kendall test, Sen's slope estimator, Bayesian procedure (Gocic and Trajkovic, 2013; Karpouzos et al., 2010) [19,20]. In this study, the linear regression method was used to analyse the precipitation trend. In this method, the trend of data may be recognised through the following equation (Gocic and Trajkovic, 2013; Ngongondo et al. 2011) [19,21].

$$Y = a + bx \dots\dots\dots (4)$$

Where x is the explanatory variable (years), Y is the dependent variable (rainfall, mm), b is the slope of the line (mm/year) and a is the intercept.

Homogeneity Test

Test of homogeneity means the detection of variation in a dataset. If such data are homogenous, this means that data were measured at the same time with the same tools and circumstances (Kang and Yusof, 2012)[22]. Two groups of homogeneity testing techniques can be

distinguished and are usually referred to as *absolute* and *relative* methods. In the first set of procedures, the statistical tests apply to each station data separately. In the second set, the test uses records from neighboring stations (named reference stations) which presumably are homogeneous. There are several techniques for testing homogeneity in the literature. The widely used techniques are the Standard Normal Homogeneity Test (SNHT), Buishand Test and Pettitt Test (Piticar and Ristoiu, 2012)[23].

The SNHT test was proposed by Alexandersson (1986)[24] to detect the inhomogeneity in the time series. In this test, null hypothesis (H_0) assumes that data are random and independent amounts that are identically and normally distributed. The alternative hypothesis (H_1) it assumes that there is a break in the mean. This test is more sensitive to breaks near the beginning and end of the time series (Ngongondo et al. 2011)[21].

In this test, a statistic T_y is used to compare the mean of first part of data (y of years) with the remaining part ($n-y$ of years) and can be written as below (Şahin and Cığızoğlu, 2010, Kang and Yusof, 2012) [25,22].

$$T_y = y\bar{z}_1 + (n - y)\bar{z}_2, \quad y = 1, 2, \dots, n \dots\dots (5)$$

$$\bar{z}_1 = \frac{1}{y} \sum_{i=1}^y \frac{(y_i - \bar{y})}{S} \dots\dots\dots (6)$$

$$\bar{z}_2 = \frac{1}{(n - y)} \sum_{i=y+1}^n \frac{(y_i - \bar{y})}{S} \dots\dots\dots (7)$$

The year y is having a break if value of T is maximum. To reject null hypothesis, the value of T_0 is compared with critical values shown in the Table 2 which depend on data size (Kang and Yusof, 2012)[22].

$$T_0 = \max T_y \dots\dots\dots (8)$$

Where n is the number of total years, S is standard deviation, \bar{y} is mean of total data and \bar{z}_1, \bar{z}_2 are the first and second data part averages.

The Buishand test was first proposed by Buishand (1982)[26] in which the null hypothesis (H_0) assumes that data are random, independent and normally distributed. The alternative hypothesis (H_1) assumes that there is a break in the mean.

This test is more sensitive to breaks in the middle of time series (Costa and Soares, 2009)[27]. The homogeneity test can be based on the cumulative deviations from the mean or adjusted partial sums, which are defined as follows (Şahin and Çiğizoğlu, 2010)[25]:

$$S_y = \sum_{i=1}^y (Y_i - \bar{Y}), \quad y = 1, 2, \dots, n \dots \dots \dots (9)$$

When data series is homogenous, the S_y values will rise and fall around zero. The year y has a break when S_y has reached a maximum (negative shift) or minimum (positive shift). Adjusted range R is obtained by

$$R = \frac{\max S_y - \min S_y}{S} \dots \dots \dots (10)$$

The value (R/\sqrt{n}) is then compared with the critical values given by Buishand (1982)[26], (See Table (2)).

The approach after Pettitt (1979)[28] is commonly applied to detect a single change-point in hydrological series or climate series with continuous data. This test is more sensitive to breaks near the middle of the time series (Costa and Soares, 2009)[27]. It is based on the rank, r_i of the y_i and ignores the normality of the series (Şahin and Çiğizoğlu, 2010; Kang and Yusuf, 2012)[25,22].

$$X_y = 2 \sum_{i=1}^y r_i - y(n + 1), \quad y_i = 1, 2, \dots, n \dots \dots (11)$$

The break occurs in year k when

$$X_k = \max |X_y| \dots \dots \dots (12)$$

The value X_k is compared with the critical value given by Pettitt (1979)[28] and shown in Table (2).

In this study, a significance level of 5% was adopted for the above-mentioned tests. If the value of the test exceeds the critical value at the specified confidence level, the null hypothesis (H_0) will be rejected, or, if the estimated value is greater than the critical value, the null hypothesis (H_0) will be rejected. While the alternative hypothesis (H_1) will be accepted at the same confidence level.

Randomness Test

Randomness in hydrological data series means that data arise out of natural reasons (McMahon and Mein, 1986)[29]. Randomness testing is summarised in calculating the data series median. Each data item is examined whether it exceeds the median. If a data item exceeds the median, then this is considered a success case (replaced by letter n_1); otherwise, it is a failure case (denoted by letter n_2). Cases that are exactly equal to the median are excluded and the number of each of n_1 and n_2 shall be counted. The value of r , which is (n_1+n_2) , shall be calculated, consequently the following equation is used to calculate the values of the standard normal variable Z (Adeloye and Montaseri, 2002)[18].

$$Z = \frac{r - \left(\frac{2(n_1 n_2)}{(n_1 + n_2)} + 1 \right)}{\sqrt{\frac{2n_1 n_2 (2n_1 n_2 - n_1 - n_2)}{(n_1 + n_2)^2 (n_1 + n_2 - 1)}}} \dots \dots \dots (13)$$

The calculated value of Z shall be compared with the normal distribution table at a confidence level of 5% with the null hypothesis (H_0) being rejected if $Z > Z_{\alpha/2}$.

Table 2. Critical values of homogeneity tests (Vezzoli et al., 2012) [30]

Test	Sig. level	Data size (n)					
		20	30	40	50	70	100
SNHT	1%	9.56	10.45	11.01	11.38	11.89	12.32
	5%	6.95	7.65	8.1	8.45	8.8	9.15
Buishand	1%	1.6	1.7	1.74	1.78	1.81	1.86
	5%	1.43	1.5	1.53	1.55	1.59	1.62
Pettitt	1%	71	133	208	239	488	841
	5%	57	107	167	235	393	677

Drought Analysis
Standardized Precipitation Index (SPI)

McKee et al. (1993)[9] developed the Standard Precipitation Index (SPI) for the purpose of defining and monitoring drought. One of the main advantages of the SPI is that it only requires rainfall data as an input, which makes it perfect for areas where data collection is not as extensive and makes its evaluation relatively easy. SPI is a standardized index, which ensures independence from geographical position as the index in question is calculated regarding the average precipitation in the same place (Cacciamani et al. 2007)[31]. SPI was developed to quantify the precipitation deficit for multiple time scales, reflecting the impact of rainfall deficiency on the availability of various water resources. They calculated the SPI for 3, 6, 12, 24, and 48-month scales to reveal the temporal performance of the impact. The SPI provides a quick and handy approach to drought analysis (Umran Komuscu, 1999)[32].

The computation of the SPI drought index is based on the long-term rainfall record (at least 30 years). The rainfall time series is fitted to a gamma distribution, which is then transformed through an equal probability transformation into a normal distribution. Positive SPI values indicate wet conditions with higher than median precipitation, and negative SPI values indicate dry conditions with lower than median precipitation (Bordi and Sutera, 2007)[33]. The procedure of calculating SPI has been detailed in McKee et al. (1993)[9] and Table 3 shows SPI drought classes.

Table 3. Drought classification based on the SPI value (McKee et al. 1993)[9]

SPI value	Drought Class
≥ 2	Extremely wet
1.5 to 1.99	Severely wet
1.0 to 1.49	Moderately wet
0 to .99	Mid wet
-.99 to 0	Mild dry
-1.49 to -1.0	Moderately dry
-1.99 to -1.5	Severely dry
-2.0≥	Extremely dry

Theory of Runs

Yevjevich (1967)[34] proposed the theory of runs as a tool for identifying drought and studying its characteristics. Theory of runs is defined as a series of similar events which are preceded and followed by different events. Theory of Runs can be used to describe the metrological drought. Drought threshold may be a fixed value in case of annual values (no seasonal variations), or may be seasonal variable values when seasonal data are used. The truncation level should be selected in each period depending on the purpose of the study. This value is often considered to be equal to the long-term *mean* or *median* for the time series to be studied. Other choices also considered such as adopting a percent of the mean or through a specific probability. The advantage of using the theory of runs in identifying drought characteristics is the ability of deriving the probabilistic aspects of drought characteristics, such as drought period length, drought accumulated deficit in an analytical concept or through using generated data when stochastic characteristics of study's basic variable are known. New methods were derived to evaluate return periods for drought events, which depend on theory of runs. Thus, this theory is ideal for studying and analysing drought risks (Cancelliere et al., 2005).

Main drought variables, which can be estimated by theory of runs and shown in Figure (4) are the drought duration, accumulated deficit and intensity (Sirdaş and Şen, 2003). Drought duration *L* is the number of successive periods experiencing rainfall deficit. The accumulated deficit *M* is the sum of rainfall deficits compared to their averages, which can be estimated by the following equation:

$$M_j = \sum_{i=1}^m \bar{X} - X_i \dots \dots \dots (14)$$

Where *M_j* is the accumulated deficit for drought *j*, *m* is number of deficits during drought *j*, *X_i* is rainfall at time *I* and \bar{X} is the average of rainfall time series. Drought intensity *I* is the ratio between the accumulated deficits to drought duration.

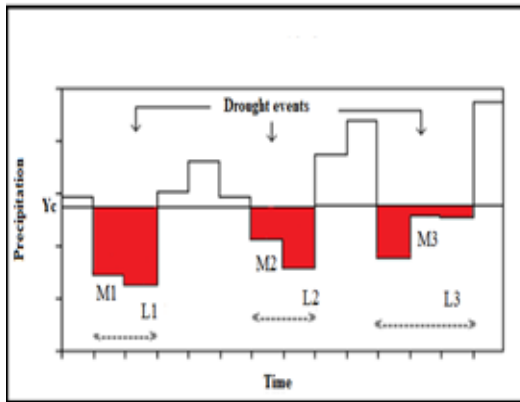


Fig. 4. Drought identification with the theory of runs

Results and Discussion

Statistical Tests of Rainfall ata

The consistency test results showed that monthly rainfall data for the selected stations within the study period experienced no breaks in rainfall data, which shows that the data of the selected stations are of consistent nature.

Regarding trend of rainfall, the trends of rainfall data series of each station were drawn to recognize if it exists. Results revealed that negative trends exist in rainfall data series of all stations covered by the study except for Hillah station as shown in the Table (4).

Table 4. Annual rainfall trend equations for the selected stations

Station	Trend equation	Station	Trend equation
Khanqin	$y^* = 361.061 - 3.1625x'$	Hai	$y = 171.688 - 1.73453x$
Qaim	$y = 134.53 - 0.617211x$	Basra	$y = 157.64 - 1.06071x$
Mosul	$y = 405.13 - 2.2086x$	Nasiriya	$y = 130.588 - 0.450202x$
Duhok	$y = 49.7657 - 1.502x$	Ramadi	$y = 139.107 - 1.16872x$
Erbil	$y = 380.181 - 1.22x$	Baghdad	$y = 158.818 - 1.9815x$
Sulaymaniyah	$y = 63.8208 - 2.01x$	Hilla	$y = 92.1434 + 1.002x$
Kirkuk	$y = 34.3646 - 1.96x$	Nukheb	$y = 76.9995 - 0.160256x$
Tel-Afer	$y = 30.0651 - 1.602x$	Baiji	$y = 219.69 - 1.01592x$
Sinjar	$y = 34.9575 - 2.102x$	Samawa	$y = 93.3741 - 4.19x$
Dukan	$y = 832.78 - 4.60x$	Hadithah	$y = 156.19 - 1.58592x$
Darbandikhan	$y = 67.2920 - 5.01x$	Rutba	$y = 135.801 - 1.02945x$

y^* : Rainfall, mm, x' : year

Rainfall data series of all stations considered in the study were tested for their homogeneity using the three methods discussed earlier. Results of Pettitt test demonstrated that all stations' data are of homogenous nature. The SNHT proved that all stations' data involved in the study are homogenous except for Ramadi station. Whereas the Buishand test showed that, all data are homogenous except for Ramadi, Haditha and Darbandikhan stations. The differences between the three tests are due to the sensitivity of each test to refraction point in the investigated rainfall data series (Costa and Soares, 2009)[27]. As the rainfall data of all stations considered in this work have been passed in one or more tests, they all are considered as successful with this test.

The randomness test results showed that the data of all stations considered in the

study are random except for Haditha and Samawah stations.

Meteorological Drought Analysis

In the current study, the SPI values were estimated for 12-months (SPI-12) for each of the considered meteorological stations individually for the years 1970-2010. Each period starts from January and ends with December. (Mckee et al. 1993; Edossa et al. 2010; Tosic et al. 2010)[9,16,37]. Figure (5) demonstrates the values of SPI-12 for stations of Dukan, Mosul, Baghdad, and Basra as examples. It could be noticed that the highest values of drought for all stations were recorded in the period 2007-2010, where SPI values declined to (-2.5) in some stations such as in Basra station which refer to an extreme drought (see Table (3)). It is worthy to mention that the highest estimated

value of SPI-12 within the study period were in Khanqin, in 1970 with a value of (-3.3). The figure shows that the severest periods of drought were within the last two decades (1997-2001) and (2007-2010) and for all stations considered in the study.

North-eastern region of Iraq is characterized by drought averages higher than the other affected areas due to high rainfall rates in the area. The largest accumulated rainfall deficit was in Darbandikhan station, with a deficit of 1769.7 mm and drought duration of 8 years, followed

by Dukan station, with a deficit of 958.5 mm, then Sulaymaniyah station with a deficit of 869.6 mm. The severest drought intensity was in Dukan station with 239.6 mm/year. The study showed that most draught events within study period were ranging between mild drought and mild wet classes according to the classification of SPI. Average percentage of dry years in the study area was 50.1%, which reveal that half of the years within the study period were experiencing droughts.

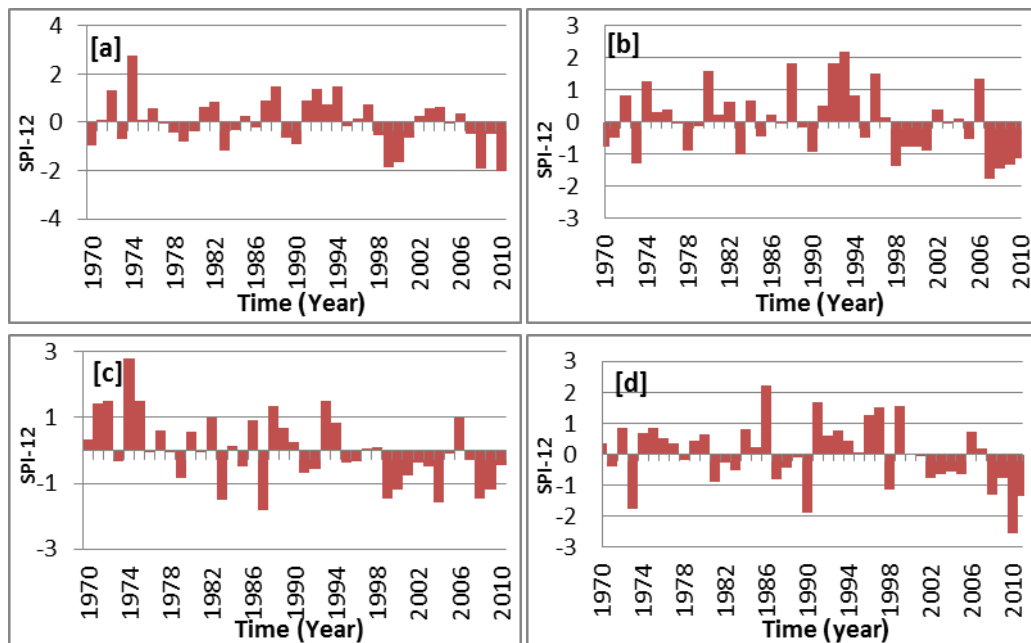


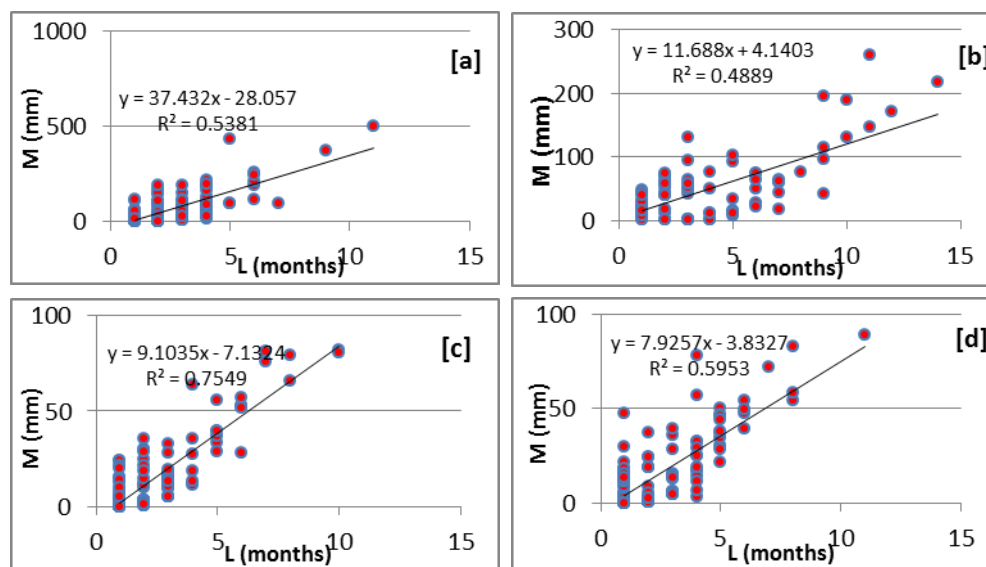
Fig. 5. SPI-12 values for a) Dukan, b) Mosul, c) Baghdad and d) Basra stations

Studying and analysing drought events will provide information that help in managing sustainable water resources in the region. Characteristics of meteorological drought were analysed, and a relationship was suggested between drought magnitude M and drought duration L for each of stations covered in the study which shown in Figure 6 and Table (5). Figure (6) shows the relation between accumulated deficit and drought duration for stations of Dukan, Mosul, Baghdad and Basra as examples. The figures reveal that the accumulated deficit value can be estimated for a certain drought duration.

Moreover, accumulated deficit increases with the increase of drought duration. This applies to all stations covered in the study with one difference that is the curve's slope from a station to another, as shown in Table (5). This relation indicates the need of water used for different purposes in critical periods and helps in managing water resources in a certain area to substitute the deficit between the available water and the demand in areas affected by drought owing to external sources by alternative water resources or through water stored during wet periods.

Table 5. Relationship between the accumulated deficit (M) and Drought duration (L) for the selected stations using SPI-12

Station	Equation	R ²	Station	Equation	R ²
Hadithah	$M = 8.7091L - 4.0326$	0.8058	Hilla	$M = 6.5534L - 3.6142$	0.6587
Nukheb	$M = 5.7344L - 3.3545$	0.7995	Samawa	$M = 4.2308L + 1.8571$	0.6352
Duhok	$M = 34.492L - 25.179$	0.7723	Rutba	$M = 4.8568L + 1.8858$	0.6216
Baghdad	$M = 9.1035L - 7.1324$	0.7549	Erbil	$M = 15.164L - 4.2128$	0.6044
Sulaymaniyah	$M = 41.463L - 28.409$	0.7295	Kirkuk	$M = 19.616L - 13.678$	0.6036
Hai	$M = 8.942L - 4.1466$	0.7128	Basra	$M = 7.9257L - 3.8327$	0.5953
Ramadi	$M = 6.5044L - 3.7938$	0.7117	Khanqin	$M = 16.307L - 10.447$	0.5831
Nasiriya	$M = 9.1074L - 4.5389$	0.6916	Sinjar	$M = 13.178L + 4.506$	0.5647
Darbandikhan	$M = 40.783L - 28.104$	0.6845	Dukan	$M = 37.432L - 28.057$	0.5381
Qaim	$M = 7.7138L - 5.2672$	0.6722	Baiji	$M = 6.584L + 4.5334$	0.5197
Tel-Afer	$M = 17.699L - 11.673$	0.6706	Mosul	$M = 11.688L + 4.1403$	0.4889

**Fig. 6.** The relationship between accumulated deficit (M) and drought duration (L) for a) Dukan, b) Mosul, c) Baghdad, and d) Basra stations using SPI-12

Conclusions

- 1- The statistical tests for the collected rainfall data showed that all rainfall data covered by the study, are consistent. Moreover, rainfall time series are of declining trend except for Hillah station and they were random in all selected stations except for Haditha and Samawah stations.
- 2- Considering homogeneity, Pettitt's test demonstrated that all stations' data are homogenous, while SNHT showed that the data selected in the study are homogenous except for Al-Ramadi station. However,

Buishand's test proved that the data are homogenous except for Al-Ramadi, Haditha and Darbandikhan stations.

- 3- Severest periods of drought in Iraq have taken place at the end of the last century and the end of the first decade of this century (1997-2001) and (2007-2010), respectively, in which the extreme drought class was dominant in most of the areas of Iraq.
- 4- The north-eastern region of Iraq is characterized by drought rates higher than the other affected areas due to high rainfall

rates in the area. The highest accumulated rainfall deficit is recorded in Darbandikhan station, with deficit depth of 1769.7 mm and drought duration of 8 years, followed by Dukan station, with depth of 958.5 mm, then Sulaymaniyah station with deficit of 869.6 mm. In addition, the highest drought intensity was in Dukan station with 239.6 mm/year.

- 5- The calculated SPI-12 for the considered stations mostly ranged between mild dry and mild wet according to SPI classification, and the percent of the dry years reached to 50.1% which means that half of the study period years were experiencing drought.

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